Recent results for the 3D Quasi-Geostrophic Equation

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Physical model

- The Quasi-Geostrophic system of equations models the evolution of the temperature in the atmosphere.
- It can be rigorously derived from the Primitive Equations (Euler equation with Coriolis force and Boussinesq approximation, see Bourgeois Beale (94) and Desjardins Grenier 98)
- At large scale, this Rosby effect is very important. Asymptotically, this leads to the so-called geostrophic balance which enforces the wind velocity to be orthogonal to the gradient of the pressure in the atmosphere (see Pedlosky).
- This model is extensively used in computations of oceanic and atmospheric circulation, for instance, to simulate global warming.



The unknown and parameters

- The dynamic is encoded in Ψ, the stream function for the geostrophic flow.
- That is, the 3D velocity (w, U) = (0, u, v) has its horizontal component verifying

$$(u, v) = (-\partial_{x_2} \Psi, \partial_{x_1} \Psi), \text{ or in short } : U = \overline{\nabla}^{\perp} \Psi,$$

where we denote

$$\overline{\nabla}\Psi=(0,\partial_{x_1}\Psi,\partial_{x_2}\Psi).$$

From the model, the buoyancy is given by

$$\Theta = \partial_z \Psi$$
.

We denote

$$\nabla_{\lambda}\phi = (\lambda \partial_{z}\phi, \partial_{x_{1}}\phi, \partial_{x_{2}}\phi), \qquad L_{\lambda}\phi = \operatorname{div}(\nabla_{\lambda}\phi).$$

where $\lambda = -1/\Theta_z^0$, is a given function, of z only, associated to the buoyancy of a reference state.

The equation

The function Ψ is solution to the following Initial Boundary value problem:

$$\begin{split} &(\partial_t + \overline{\nabla}^\perp \Psi \cdot \nabla)(L_\lambda \Psi + \beta_0 x_2) = 0, \qquad t > 0, \quad z > 0, \quad x \in \mathbb{R}^2, \\ &(\partial_t + \overline{\nabla}^\perp \Psi \cdot \nabla)\gamma_\nu(\nabla_\lambda \Psi) = \nu \overline{\Delta} \Psi, \qquad t > 0, \quad z = 0, \quad x \in \mathbb{R}^2, \\ &\Psi(0, z, x) = \Psi^0(z, x). \qquad t = 0, \quad z > 0, \quad x \in \mathbb{R}^2. \end{split}$$

The parameter β_0 comes from the usual β -plane approximation. The term $\gamma_{\nu}(\nabla_{\lambda}\Psi)$ stands for the Neumann condition at z=0 associated to the operator $L_{\lambda}\Psi$. If λ is regular, this coincides with $-\lambda(0)\partial_z\Psi(0,\cdot)$. The ν term is due to the Eckman pumping. $\nu=0$ corresponds to the inviscid case.

- Both, the value of the elliptic operator $L_{\lambda}\Psi$, and the Neumann condition $\gamma_{\nu}(\nabla_{\lambda}\Psi)$ at the boundary z=0, are advected by the stratified flow with velocity $U=\overline{\nabla}^{\perp}\Psi$. At each time, Ψ can be recovered, solving the boundary value elliptic equation.
- Main difficulty: treatment of the boundary condition.

The inviscid case

We assume that $\nu = 0$, and that there exists $\Lambda > 0$ such that

$$\frac{1}{\Lambda} \le \lambda(z) \le \Lambda, \quad \text{for } z \in \mathbb{R}^+.$$

Theorem (Puel-V.)

Consider an initial value Ψ^0 such that

$$\mathcal{L}_{\lambda}\Psi^{0}$$
 and $\nabla_{\lambda}\Psi^{0}$ are in $\mathcal{L}^{2}(\mathbb{R}^{+}\times\mathbb{R}^{2}), \qquad \gamma_{\nu}(\nabla_{\lambda}\Psi^{0}) \in \mathcal{L}^{2}(\mathbb{R}^{2}).$

Then, there exists Ψ weak solution to the Quasi-Geostrophic equation on $(0,\infty)\times\mathbb{R}^+\times\mathbb{R}^2$, such that for every T>0, $\nabla_\lambda\Psi\in L^\infty(0,T;L^2(\mathbb{R}^+\times\mathbb{R}^2))\cap C^0(0,T;L^2_{loc}(\mathbb{R}^+\times\mathbb{R}^2))$.



The case with Eckman pumping

We assume that $\lambda(z) = 1$, and $\nu > 0$.

Theorem (Novack-V.)

Consider an initial value $\nabla \Psi^0 \in L^2(\mathbb{R}^3_+) \cap H^p((0,\infty) \times \mathbb{R}^2)$ with $p \geq 3$. Then, there exists a unique global solution $\nabla \Psi$ to the Quasigeostrophic equation on $(0,\infty) \times \mathbb{R}^+ \times \mathbb{R}^2$, such that for every T>0, $\nabla_\lambda \Psi \in C^0(0,T;H^p(\mathbb{R}^+ \times \mathbb{R}^2))$.

Especially, if the initial is smooth (C^{∞}), then the unique solution is also smooth.

Main difficulty

To simplify the exposition, let us consider the case with out forcing with $\beta=0$, and $\lambda=1$.

$$\begin{split} &(\partial_t + \overline{\nabla}^\perp \Psi \cdot \nabla)(\Delta \Psi) = 0, & \text{for } z > 0, \\ &(\partial_t + \overline{\nabla}^\perp \Psi \cdot \nabla)(\partial_z \Psi) = 0, & \text{for } z = 0, \\ &\Psi(0, z, x) = \Psi^0(z, x). & t = 0. \end{split}$$

• A priori estimates: for any $1 \le p \le \infty$:

$$\begin{split} \|\Delta\Psi(t)\|_{L^{p}(\mathbb{R}^{+}\times\mathbb{R}^{2})} &\leq \|\Delta\Psi(0)\|_{L^{p}(\mathbb{R}^{+}\times\mathbb{R}^{2})}, \\ \|\partial_{z}\Psi(t,0)\|_{L^{p}(\mathbb{R}^{2})} &\leq \|\partial_{z}\Psi(0,0)\|_{L^{p}(\times\mathbb{R}^{2})}, \end{split}$$

• No compactness on the trace of $\partial_z \Psi$ at z = 0!

A special case: the Surface Quasi Geostrophic Equation

- If $\Delta \Psi(0) = 0$, then $\Delta \Psi(t) = 0$ for all $t \ge 0$.
- Denote $\theta = \partial_z \Psi$ defined at z = 0. Then θ is solution to

$$\partial_t \theta + U \cdot \nabla \theta = 0, \qquad t > 0, (x, y) \in \mathbb{R}^2,$$
 (1)

$$\theta = \theta_0, \qquad t = 0, (x, y) \in \mathbb{R}^2,$$
 (2)

and the velocity U can be expressed in \mathbb{R}^2 , via a nonlocal operator, as

$$U = \nabla^{\perp} \Delta^{-1/2} \theta.$$

- This model has been popularized as a toy problem for 3D fluid mechanics (see Constantin, Majda, Held, Pierrehumbert, Garner, Swanson ...).
- Our theorem extends to QG the result of Tabak for SQG, using different techniques.

A new formulation (1)

- The proof does NOT use (and does not show) compactness on the trace of $\partial_z \Psi$ at z = 0.
- It is based on a reformulation of the problem into a system of equations (without equation on the trace).
- The stability (and compactness) for this problem is pretty simple.
- We then have to show the equivalence between the two formulations.

A new formulation (2)

• Consider the Hodge decomposition in $L^2(\mathbb{R}^+ \times \mathbb{R}^2)$:

$$u = \nabla_{\lambda} \phi + \operatorname{curl} v = \mathbb{P}_{\lambda} u + \mathbb{P}_{\operatorname{curl}} u,$$

with $\operatorname{curl} v \cdot \nu = 0$ at z = 0.

The QG problem can be reformulated as

$$\partial_t \nabla_\lambda \Psi + \mathbb{P}_\lambda (\bar{\nabla} \Psi^\perp \cdot \bar{\nabla} \nabla_\lambda \Psi) = 0, \quad \text{on } \mathbb{R}^+ \times \mathbb{R}^2 \times \mathbb{R}^+.$$

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Taking the div of the equation gives the first QG equation, thanks to

$$\operatorname{div}(\mathbb{P}_{\lambda}\cdot)=\operatorname{div}(\cdot), \qquad \partial_{i}(\bar{\nabla}\Psi)^{\perp}\cdot\bar{\nabla}\partial_{i}\Psi=0.$$

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 Taking the trace of the system a z = 0 gives (formally) the trace condition of QG, since formally, at z = 0

$$\mathbb{P}_{\lambda}(f) \cdot \nu = f \cdot \nu.$$



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Note that we have $\mathbb{P}_{\text{Curl}}(\nabla_{\lambda}\Psi)=0$.

• Euler Equation:

$$\partial_t \text{curl} v + \mathbb{P}_{\text{curl}}[\text{curl} v \cdot \nabla \text{curl} v] = 0, \qquad (t, x, z) \in \mathbb{R}^+ \times \mathbb{R}^2 \times \mathbb{R}^+.$$

with
$$\mathbb{P}_{\lambda}(\operatorname{curl} v) = 0$$
 (that is $\operatorname{curl} v \cdot \nu = 0$ at $z = 0$).

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- QG equation:

$$\partial_t \nabla_\lambda \Psi + \mathbb{P}_\lambda (\bar{\nabla} \Psi^\perp \cdot \bar{\nabla} \nabla_\lambda \Psi) = 0, \quad \text{on } \mathbb{R}^+ \times \mathbb{R}^2 \times \mathbb{R}^+.$$

Note that we have $\mathbb{P}_{\text{curl}}(\nabla_{\lambda}\Psi)=0$.

• Euler Equation:

$$\partial_t \text{curl} \boldsymbol{v} + \mathbb{P}_{\text{Curl}}[\text{curl} \boldsymbol{v} \cdot \nabla \text{curl} \boldsymbol{v}] = 0, \qquad (t, \boldsymbol{x}, \boldsymbol{z}) \in \mathbb{R}^+ \times \mathbb{R}^2 \times \mathbb{R}^+.$$

with
$$\mathbb{P}_{\lambda}(\text{curl} v) = 0$$
 (that is $\text{curl} v \cdot v = 0$ at $z = 0$).

- The first equation of QG is equivalent to the vorticity equation of Euler:
 - QG:

$$\partial_t \operatorname{\mathsf{div}}
abla_\lambda \Psi + \bar{
abla} \Psi^\perp \cdot \bar{
abla} (\operatorname{\mathsf{div}}
abla_\lambda \Psi) = 0$$

• Euler:

$$\partial_t \operatorname{curlcurl} v + \operatorname{curl} v \cdot \nabla (\operatorname{curlcurl} v) = 0.$$

Proof of the Theorem

Compactness holds for the reformulated problem.

Note that \mathbb{P}_{λ} commutes with $\bar{\nabla}$, and is continuous in L^{p} .

The two formulation of QG are equivalent.

A special case: the Surface Quasi Geostrophic Equation

- If $\Delta \Psi(0) = 0$, then $\Delta \Psi(t) = 0$ for all $t \ge 0$.
- Denote $\theta = \partial_z \Psi$ defined at z = 0. Then θ is solution to

$$\partial_t \theta + U \cdot \nabla \theta = \nu \overline{\Delta} \Psi, \qquad t > 0, (x, y) \in \mathbb{R}^2,$$
 (3)

$$\theta = \theta_0, \qquad t = 0, (x, y) \in \mathbb{R}^2,$$
 (4)

and the velocity U and the Eckman pumping term $\nu \overline{\Delta} \Psi$ can be expressed in \mathbb{R}^2 , via a nonlocal operator, as

$$U = \nabla^{\perp} \Delta^{-1/2} \theta, \qquad \nu \overline{\Delta} \Psi = \nu \Delta^{1/2} \theta.$$

• The propagation of regularity for this equation has first been proved by Kiselev, Nazarov and Volberg. The global regularity of solutions with initial values in L² has been proved first by Caffarelli V. Several other proofs has been proposed by Kiselev and Volberg, and Constantin and Vicol.

The 3D case

- In the 3D case, the equation in z > 0 is hyperbolic. We can have only propagation of regularity.
- We need the propagation of almost Lipschitz norm (possible log Lipschitz).
- The regularization effects on the boundary are only C^{α} .

Sketch of the proof (1)

We decompose the solution $\Psi=\Psi_1+\Psi_2$ into two components as follows:

$$\begin{cases} \Delta \Psi_1 = 0 \\ \partial_\nu \Psi_1 = \partial_\nu \Psi \end{cases} \qquad \begin{cases} \Delta \Psi_2 = \Delta \Psi \\ \partial_\nu \Psi_2 = 0. \end{cases}$$

 The bulk of the proof is centered around verifying a version of the Beale-Kato-Majda criterion.

Sketch of the proof (1)

We decompose the solution $\Psi=\Psi_1+\Psi_2$ into two components as follows:

$$\begin{cases} \Delta \Psi_1 = 0 \\ \partial_\nu \Psi_1 = \partial_\nu \Psi \end{cases} \qquad \begin{cases} \Delta \Psi_2 = \Delta \Psi \\ \partial_\nu \Psi_2 = 0. \end{cases}$$

- The bulk of the proof is centered around verifying a version of the Beale-Kato-Majda criterion.
- The equation on the boundary of $\theta = \partial_{\nu} \Psi_1$ is of the form

$$\partial_t \theta + \mathbf{u} \cdot \overline{\nabla} \theta + (-\overline{\Delta})^{\frac{1}{2}} \theta = \mathbf{f},$$

with $f = \overline{\Delta} \Psi_2$.

- The natural a priori bound for f is in $B_{\infty,\infty}^0$.
- Using De Giorgi techniques, we get θ bounded in C^{α} .

Sketch of the proof (2)

A careful potential theory argument shows that solutions to

$$\partial_t g + (-\overline{\Delta})^{\frac{1}{2}} \theta = h \in L^{\infty}(0, T; B^0_{\infty,\infty})$$

are bounded in $L^{\infty}(0, T; B^{1}_{\infty,\infty})$. This is because the operator $(-\overline{\Delta})^{\frac{1}{2}}$ is of order 1.

- Bootstrapping an increase of regularity on the C^{α} on the drift-diffusion equation on the boundary gives that $\partial_{\nu}\Psi \in L^{\infty}(0,T;\mathcal{B}^{1}_{\infty,\infty})$ on the boundary.
- Using that the flow is stratified, this gives the "almost Lipschitz" bound needed on the velocity in z > 0 generated by the boundary.

Thank you

Thank you!!